

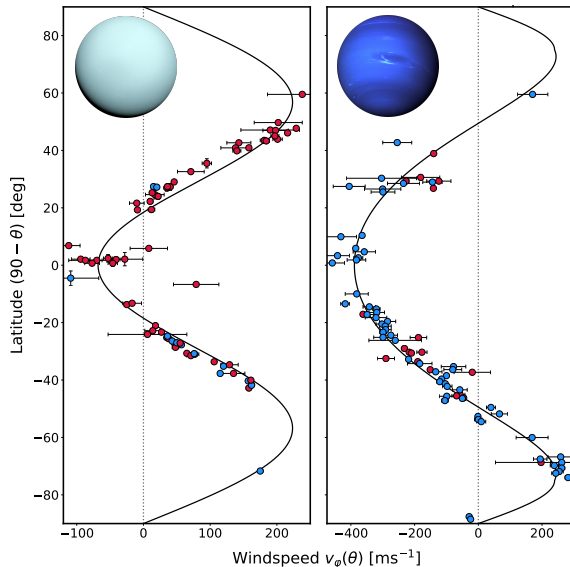
Constraining the Depth of the Winds on **Uranus** and **Neptune** via Ohmic Dissipation

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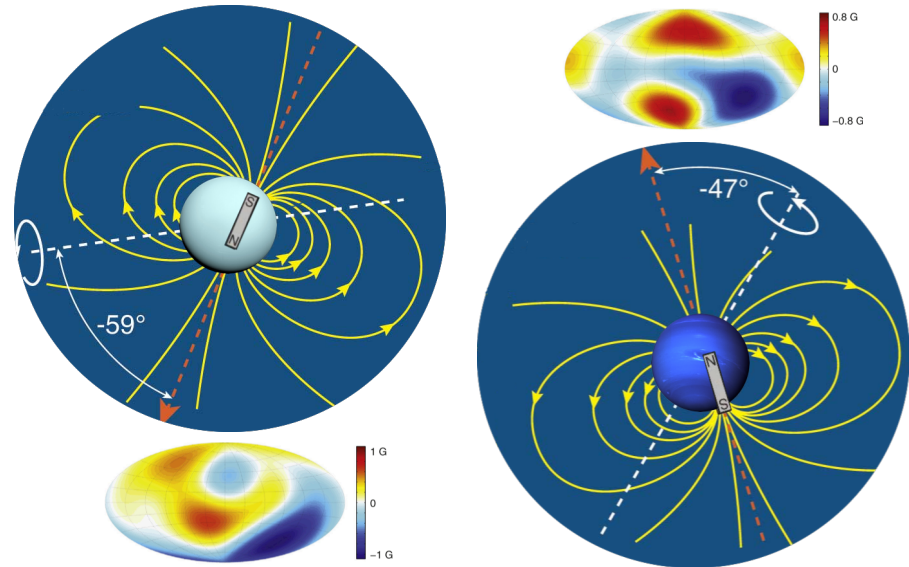
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Equatorially symmetric
and **strong** surface winds.

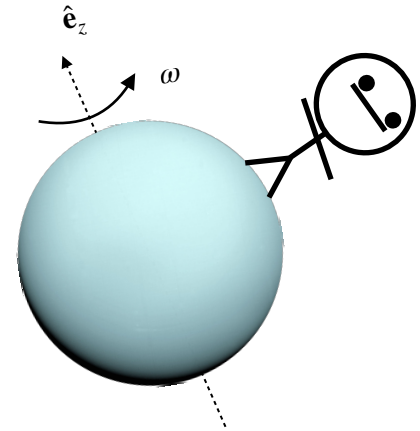
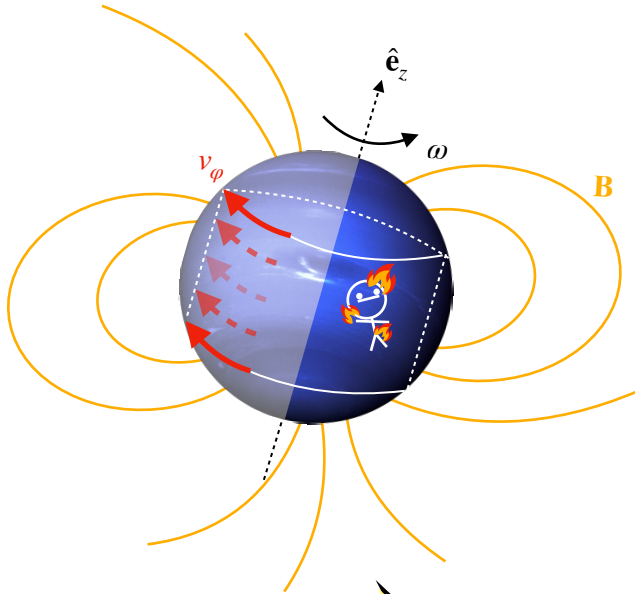


Non-axisymmetric and
multipolar magnetic fields.



Fast planetary rotation imposes geostrophic flow.
 → *weak variation of wind velocities along the z-direction*

Winds penetrate inside the planets along cylinders parallel to the rotation axis.



Winds interact with **the magnetic field** and induce **a current!**

Currents lead to **Ohmic dissipation!**

Current density: $\mathbf{j} = \sigma(\mathbf{E} + \mathbf{U} \times \mathbf{B})$

Total built-up Ohmic dissipation: $P_{\text{tot}} = \int_V \frac{\mathbf{j}^2}{\sigma} dV$

⚡ How do you estimate the conductivity? ⚡

🔥 How much dissipation is too much? 🔥

Estimating electrical conductivity

- Assume a mixture of **hydrogen**, **helium** and heavier elements.
- Represent heavy elements with **water**.
- The mixture starts to conduct **ionically** with depth.

Relevant species: $\beta \in \{H_2O, HO^-, O^{2-}, H_2, H^+\}$
 under conditions of interest

► **Does not actually have +1 charge!**

Detailed description in:
[doi:10.1093/mnras/staa2461](https://doi.org/10.1093/mnras/staa2461)

Ionic conductivity

via the fluctuation-dissipation theorem
 assuming the charges are time-independent

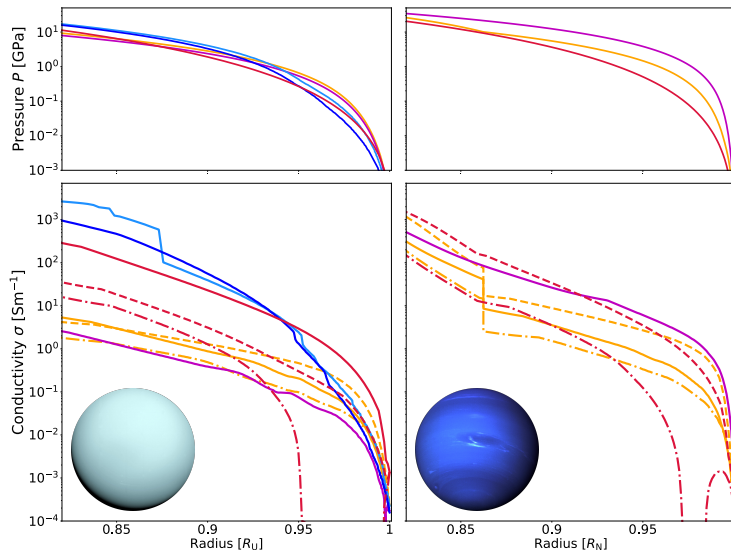
$$\sigma_{\text{ion}} = \underbrace{\sum_{\beta} q_{\beta}^2 \frac{\xi_{H,\beta} n_H}{k_B T} D_{\beta}}_{\text{Summed for hydrogen}} + \underbrace{\sum_{\beta} q_{\beta}^2 \frac{\xi_{O,\beta} n_O}{k_B T} D_{\beta}}_{\text{Summed for oxygen}}$$

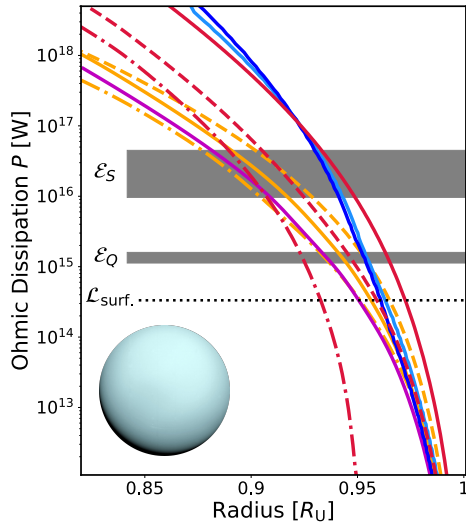
Labels for the equation:

- Charge of β (points to q_{β})
- Fraction of H in β (points to $\xi_{H,\beta}$)
- Number density of H (points to n_H)
- Diffusion coefficient of β (points to D_{β})

Ionic conductivity

of various structure models in literature
 calculated with our prescription





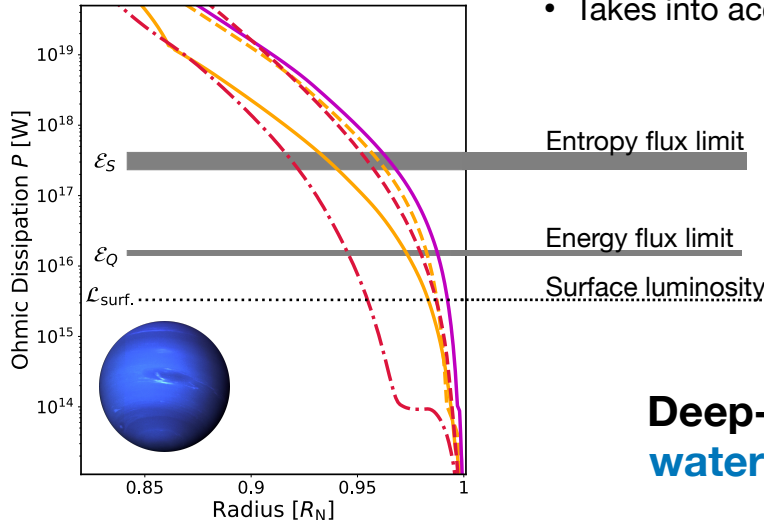
Energy flux budget:
 assuming adiabatic cooling balances
 dissipative heating at each layer

$$P_{\text{tot}} \lesssim \mathcal{E}_Q = \int_{r_i}^R \frac{Q_A}{-T/(\partial T/\partial r)} r^2 dr$$

- Takes into account the dissipative heating throughout the interior.
- Assumes advection Q_A is the major contributor to the heat flux.

Entropy flux budget: $P_{\text{tot}} \lesssim \mathcal{E}_S = \frac{T(r)}{T_0} L_{\text{surf}}$

- Takes into account the entropy generation throughout the interior.



The limits are violated above $0.90 R_{U,N}$.

**Deep-seated winds are unlikely if
 water is present in outer regions!**